

# Kriging-based Interpolation of Dissolved Oxygen in Chesapeake Bay

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CBP Tidal Monitoring and Analysis Workgroup Dissolved Oxygen Seminar

May 7, 2013

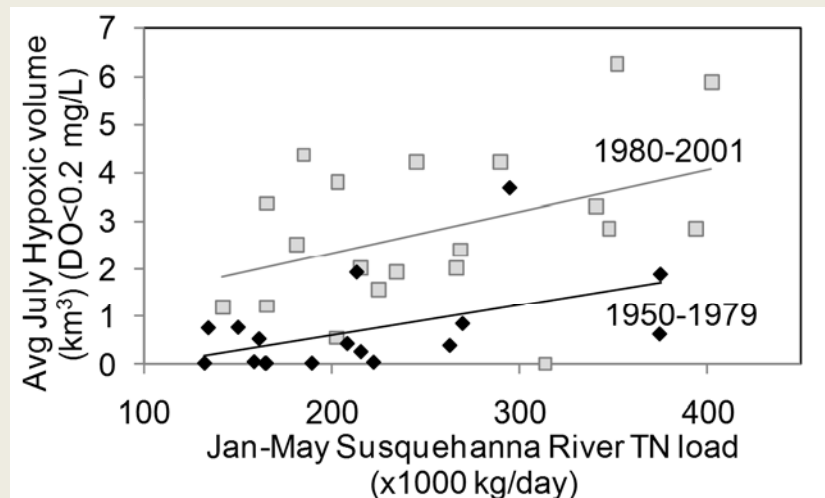
# Guiding Questions/Outline

- What method(s) do you use for assessing DO?
- Why did you choose that methodology?
- What have you found drives DO patterns in the Bay?
- What lessons have you learned that we should consider as we move forward?

# Outline

- What method(s) do you use for assessing DO?
- Why did you choose that methodology?
  - For analysis of the relationship between hypoxic volume and nitrogen loads in Chesapeake Bay
- What have you found drives DO patterns in the Bay?
- What lessons have you learned that we should consider as we move forward?

# Hypoxia: Shift in relation to nutrient loads?



Based on: Hagy et al. (2004)  
*Estuaries* 27: 634-658.  
and Kemp et al. (2005)  
*Marine Ecology Progress Series* 303: 1-29.

investigate

**CBEO**

Chesapeake Bay Environmental Observatory

CBEO website: <http://cbeo.communitymodeling.org>

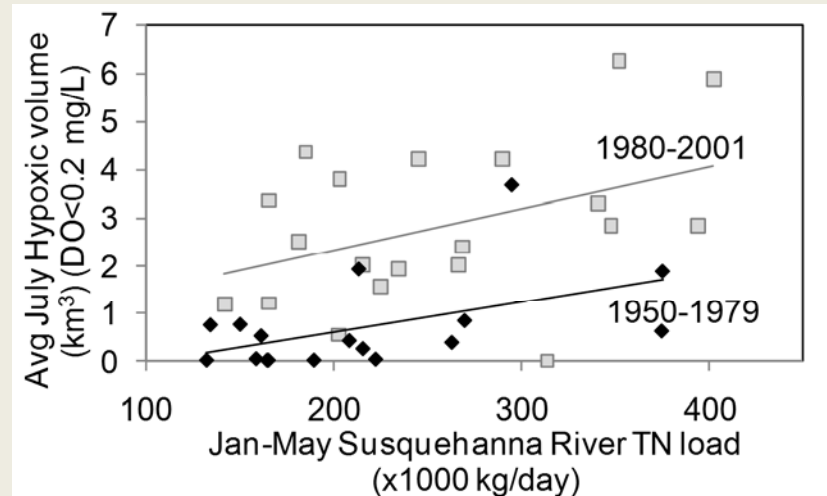
## Data Sets

CBP data  
USGS River Monitoring  
Historic Chesapeake Bay  
Institute data  
Model outputs, etc...



New  
Tools/  
Methods

# Hypoxia: Shift in relation to nutrient loads?



Based on: Hagy et al. (2004)  
*Estuaries* 27: 634-658.  
and Kemp et al. (2005)  
*Marine Ecology Progress Series* 303: 1-29.

## Hypotheses investigated by CBEO team

- Is the shift an artifact of the interpolation method used?
- Has there been a change in the stratification of the bay?
- Has there been an increase in other nutrient loads besides TN from the Susquehanna River?
- Could enhanced nutrient recycling from the sediment account for the difference?

# Outline

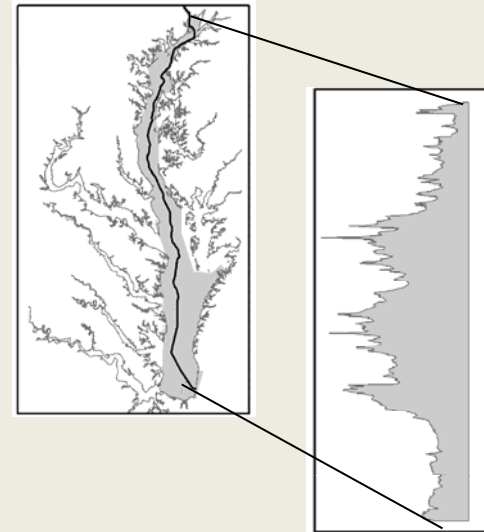
- What method(s) do you use for assessing DO?

- DO analysis: Hypoxic volume
- Spatial interpolation: Kriging
- Extent:

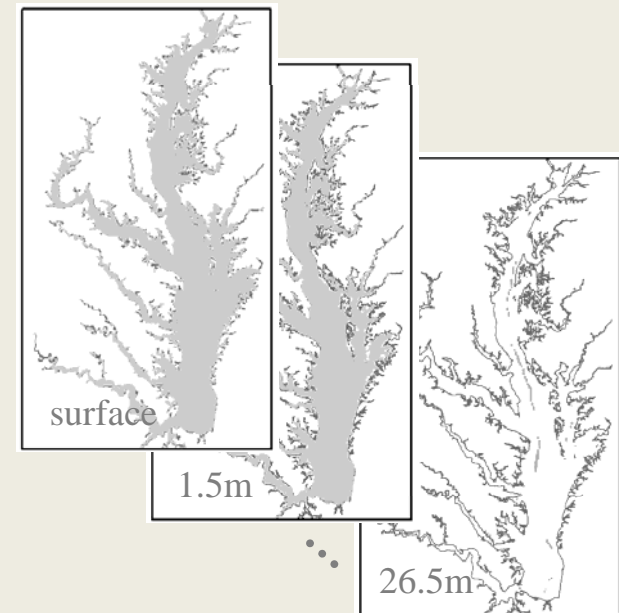
- Main-stem only
- Full-bay

- Why did you choose that methodology?
- What have you found drives DO patterns in the Bay?
- What lessons have you learned that we should consider as we move forward?

## 1) Main-stem method



## 2) Full-bay method



# Spatial Interpolation: Different methods

General: Weighted average prediction

$$\hat{Y}(s_0) = \sum_{i=1}^n w(s_i)Y(s_i)$$

$s_0$  = location (x,y) to predict at

$s_i$  = different x,y location with measured value

$\hat{Y}(s_0)$  = value predicted at location  $s_0$

$Y(s_i)$  = value measured at location  $s_i$

$w(s_i)$  = weight assigned to measured value at  $s_i$

Non-statistical

Statistical

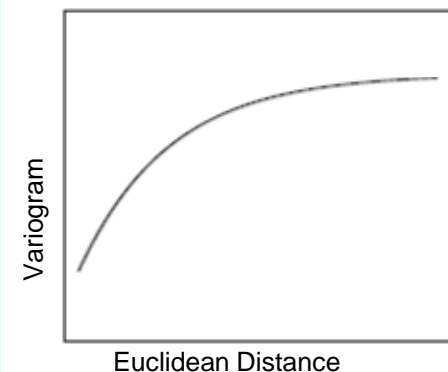
Inverse distance weighting (IDW)  
squared predictor weights

$$w(s_i) = \frac{1/d_i^2}{\sum_{j=1}^n (1/d_j^2)}$$

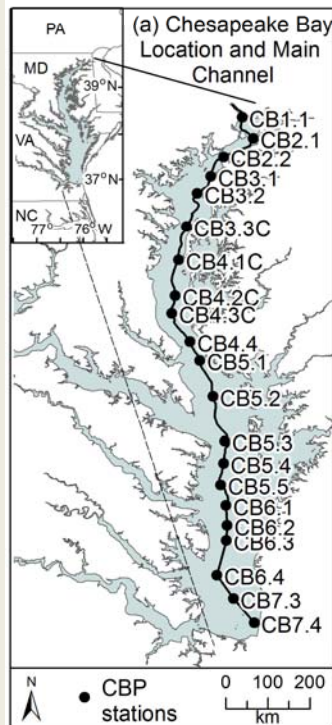
$d_i$  = distance between  $s_0$  and location  $s_i$

Kriging predictor weights

$w(s_i)$  is dependent on the spatial relationship between the observed data, as quantified with a variogram

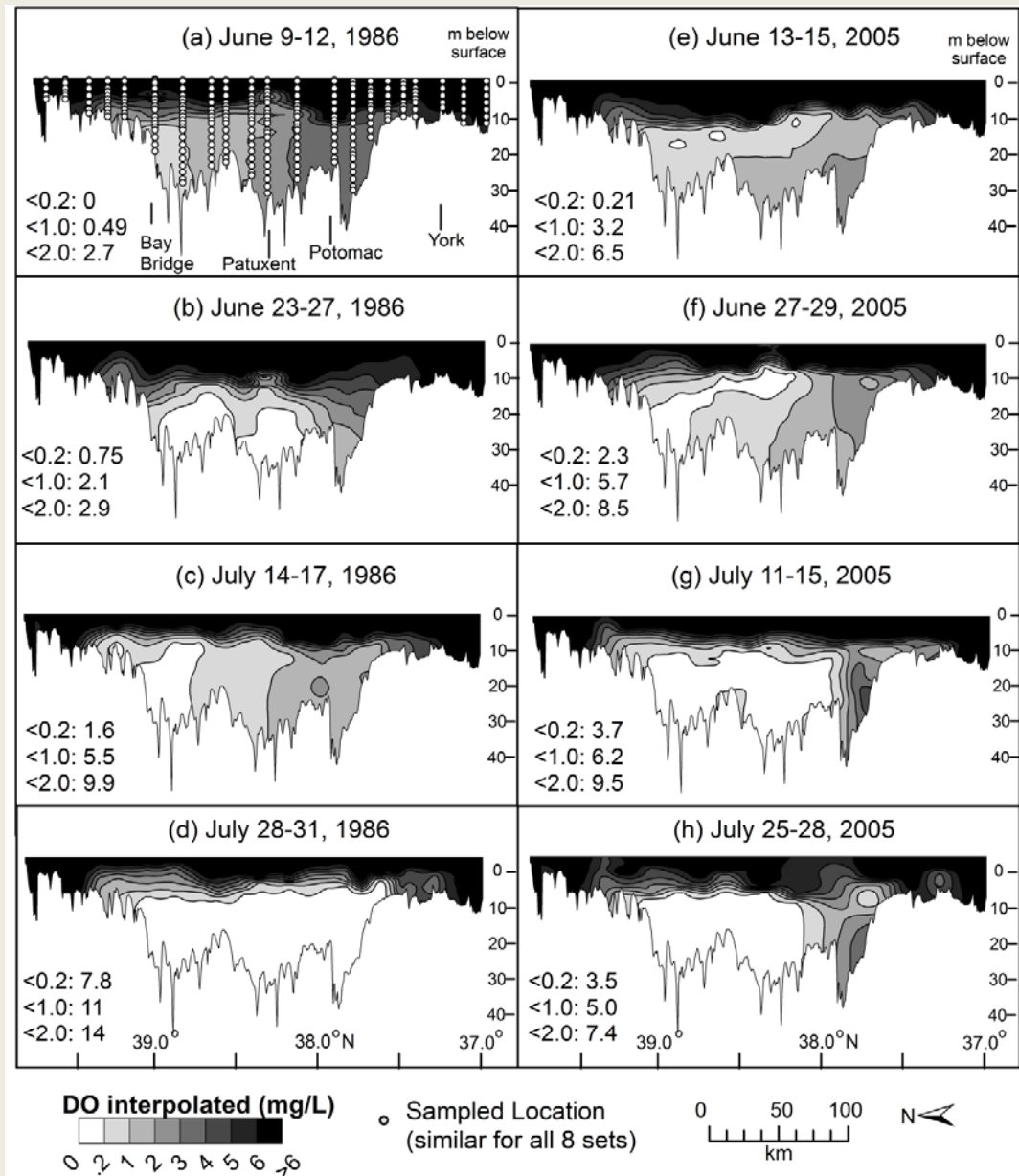


# Method 1: Main stem DO



Murphy, Kemp, and Ball . 2011.  
*Estuaries and Coasts* 34:1293-1309.

## Method 1: Main Channel

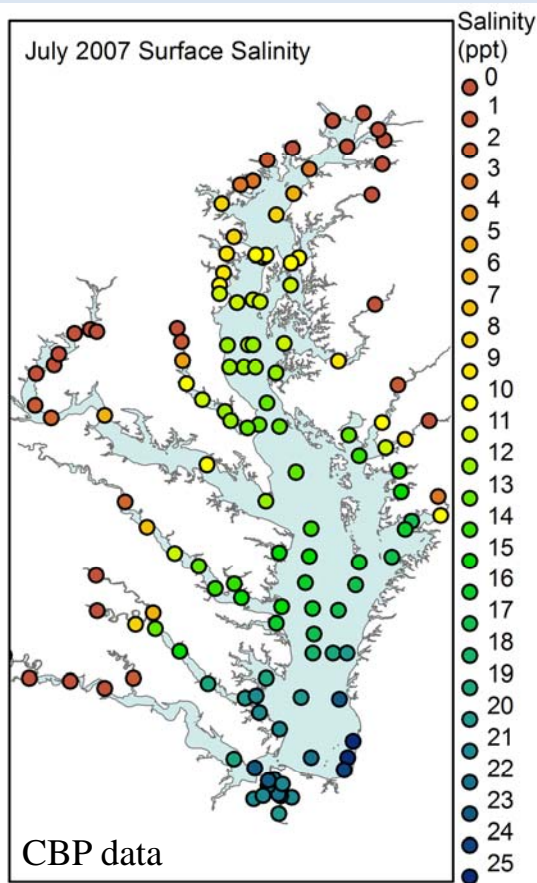




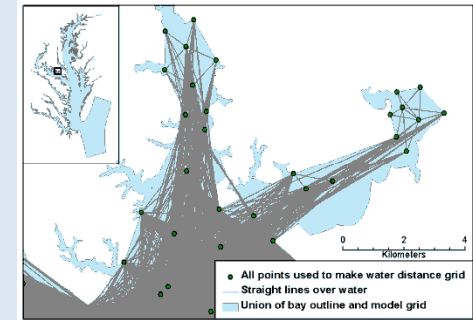


# Method 2: Full-bay method – surface only

Use all data below fall-line in Bay



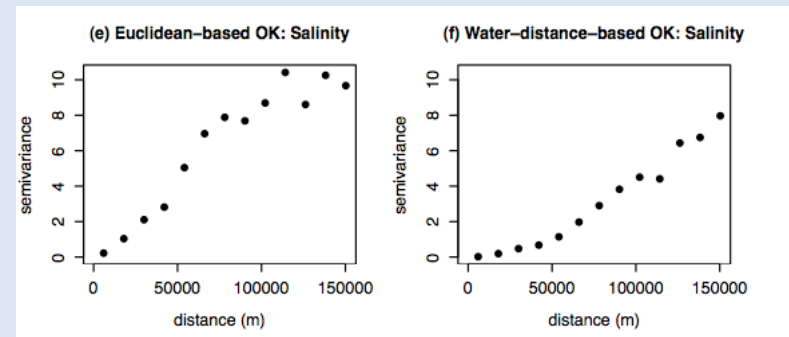
Compute “water distance” between all samples and points on a prediction grid



Transform matrix of water distances into Euclidean distances with multi-dimensional scaling

$$\begin{bmatrix} d_{11} & \dots & d_{n1} \\ \dots & \dots & \dots \\ d_{1n} & \dots & d_{nn} \end{bmatrix}$$

Use transformed distances to compute variograms and perform kriging

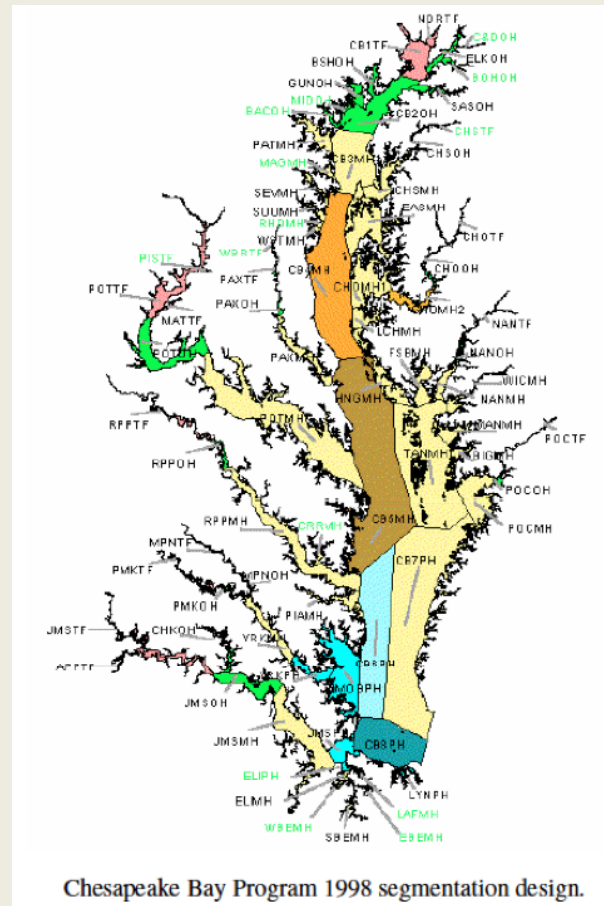


## Compare to two other methods

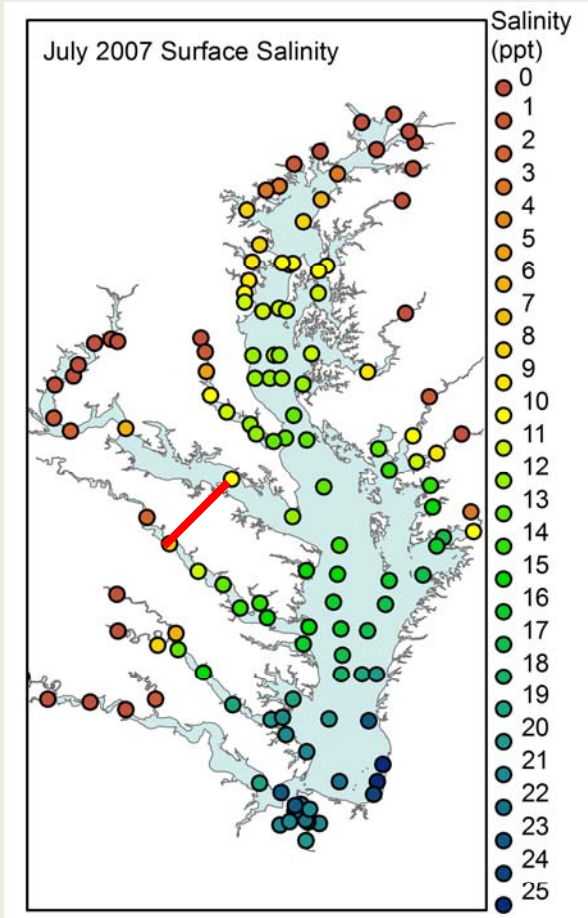
# Perform comparisons for:

- Seven parameters in surface layer
- Multiple datasets in 2007 and 2008

### a) Segmented IDW



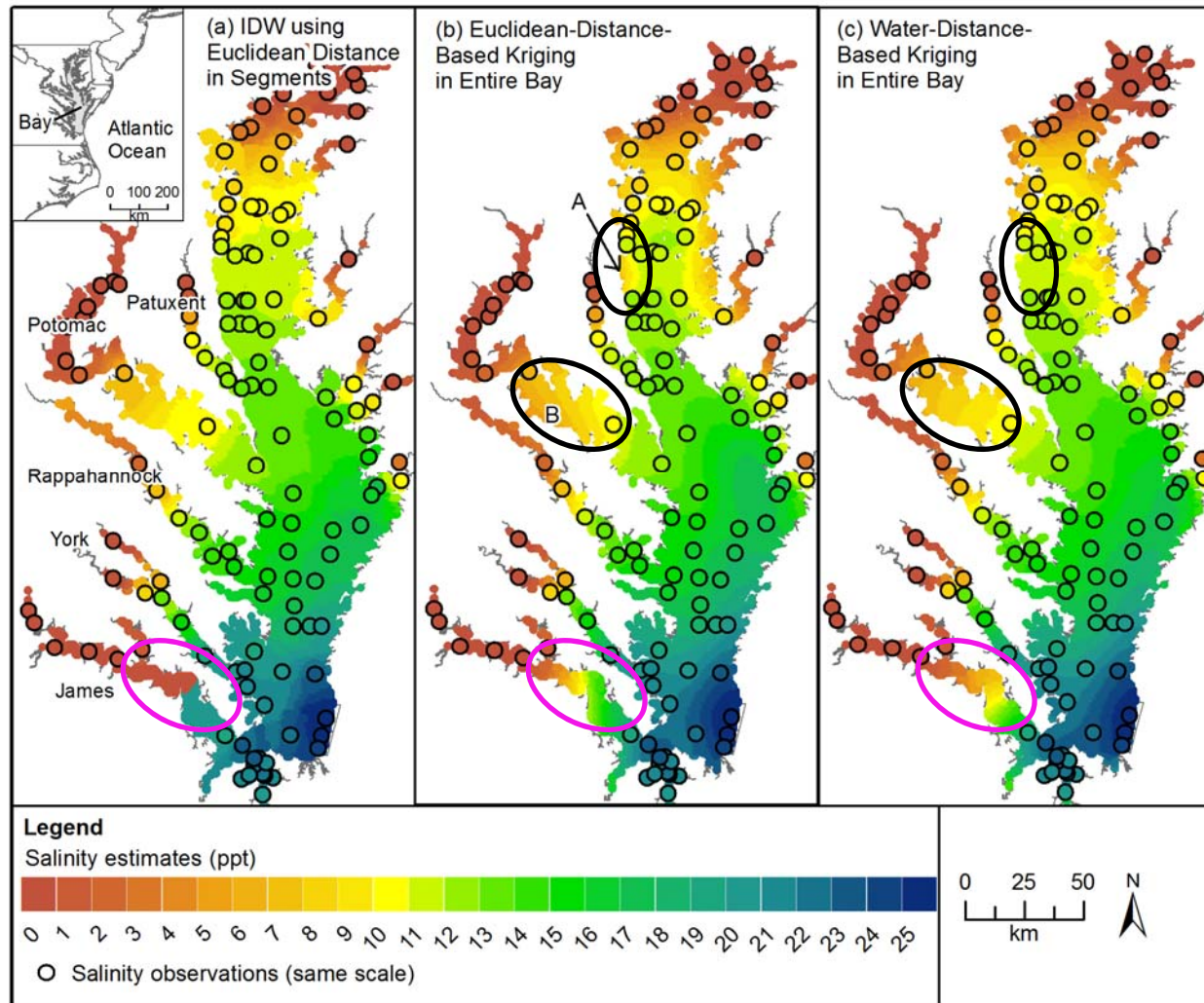
### b) Kriging with Euclidean distance





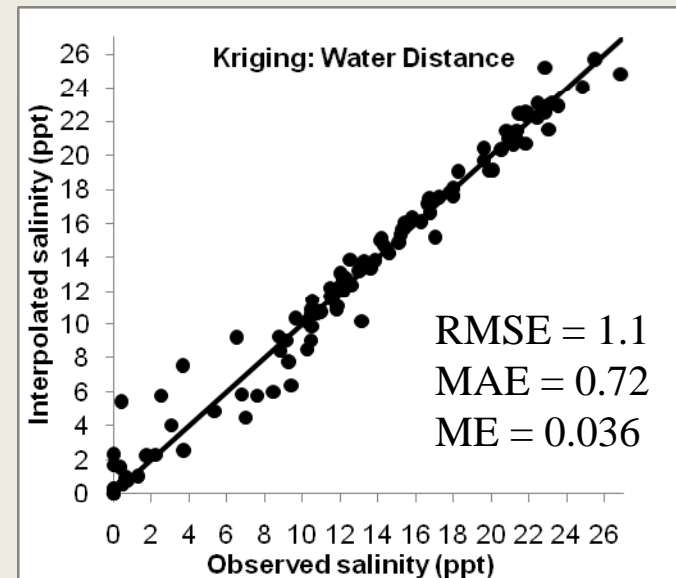
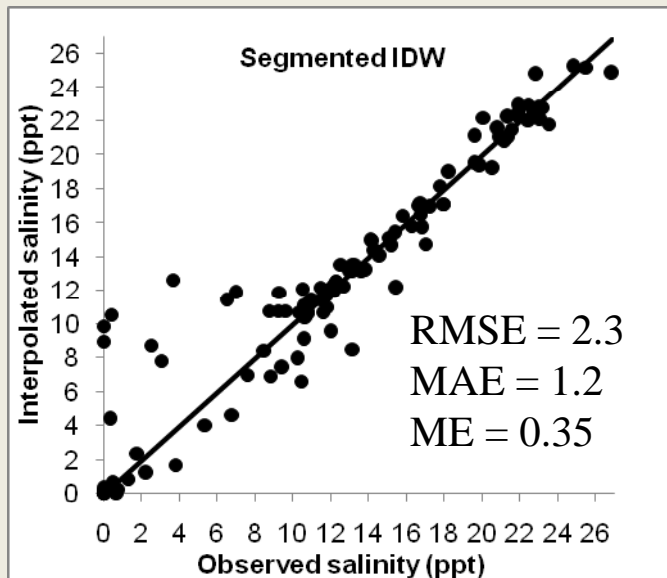
# Surface interpolation comparisons

## July 2007 Surface Salinity



## Cross validation of surface interpolation results

- Hold out each sample, one at a time, and predict at the sample location
- Examine on plots, compute average errors



## Water-distance-based method: Layers

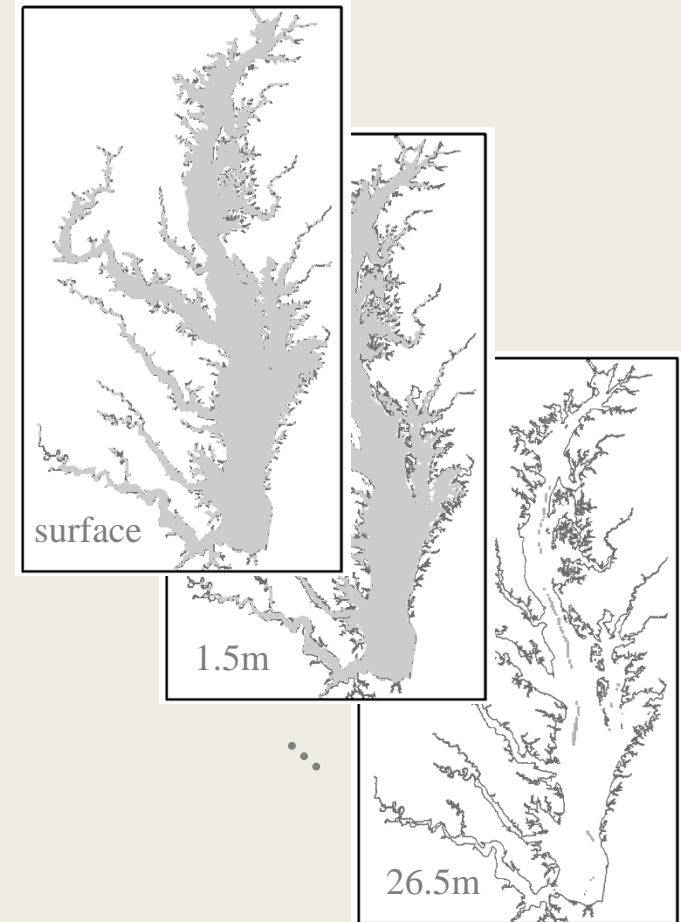
- Apply water-distance-based interpolation to data at every 1-m of depth
- To improve the vertical estimates -- use an indicator covariate as to whether each depth is above or below the pycnocline

$$DO(s) = \beta_0 + \beta_1 P(s) + \varepsilon(s)$$

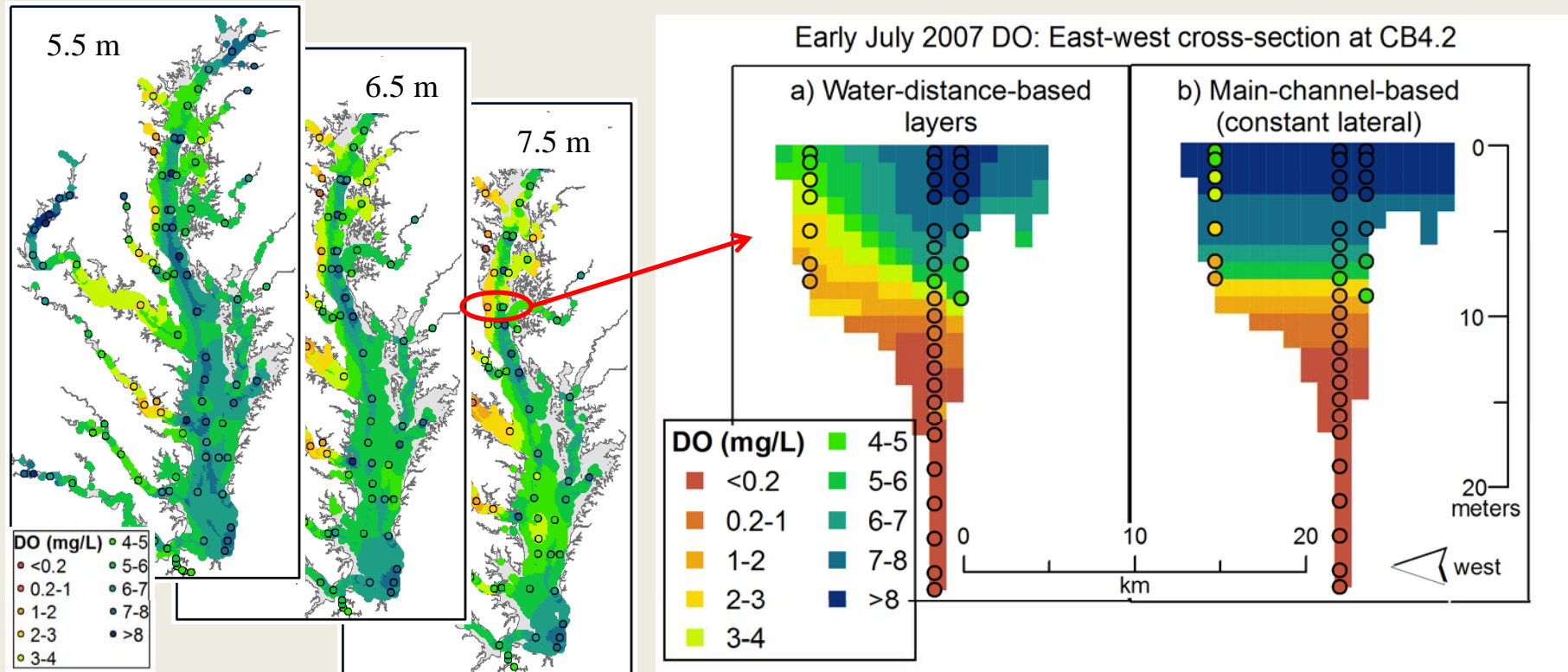
$P(s)$  = pycnocline indicator

$\beta_1$  = fitted coefficient

$\varepsilon(s)$  = spatially varying random error



## Water-distance-based method: 3-D output



## Overall: Water-distance-based method

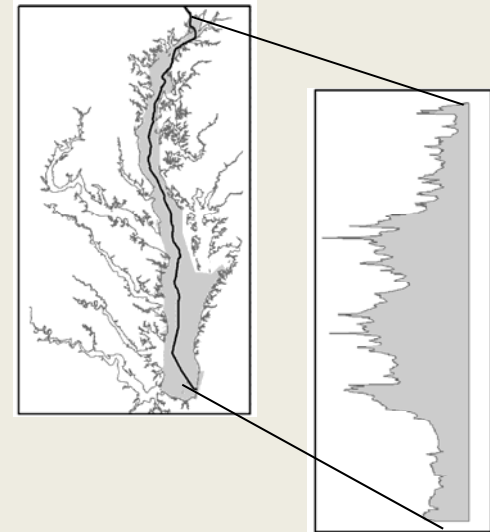
- Good in shallow water
- Captures lateral mainstem patterns
- Some problems at deeper depths
- Can be used to compute hypoxic volume for the full bay



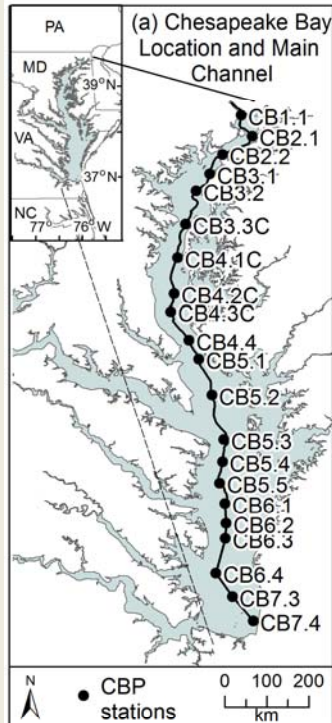
# Outline

- What method(s) do you use for assessing DO?
- Why did you choose that methodology?
- What have you found drives DO patterns in the Bay?
  - Averaging period impacts perceived trend
  - Stratification and nitrogen loads
- What lessons have you learned that we should consider as we move forward?

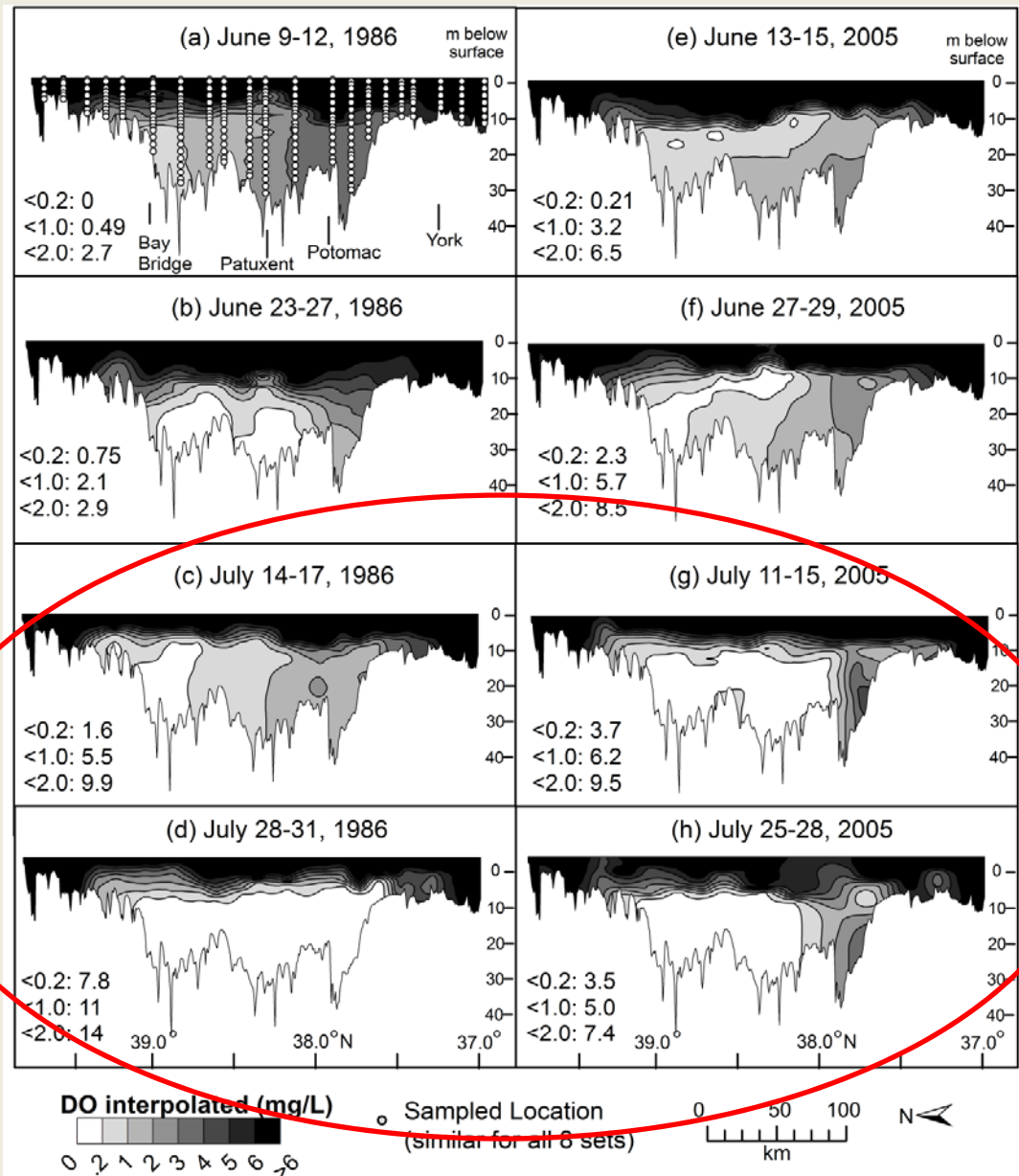
*Main-stem method*



# Main channel DO examples: June and July 1986 and 2005

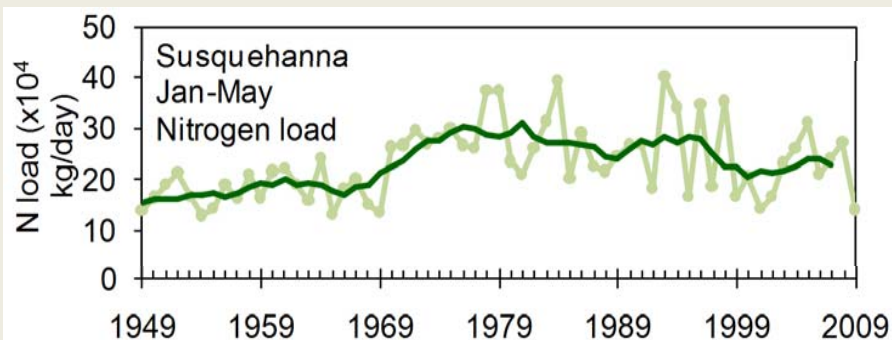
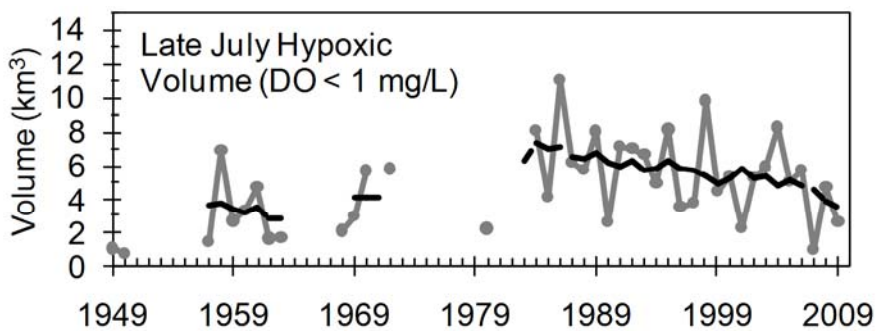
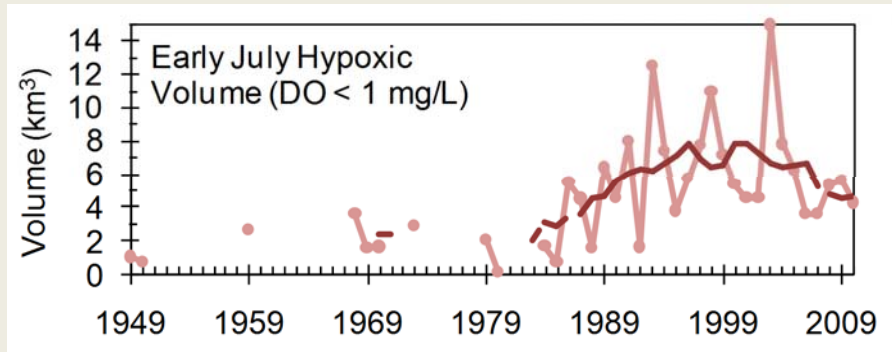


## Results Part 1: Main-channel hypoxia



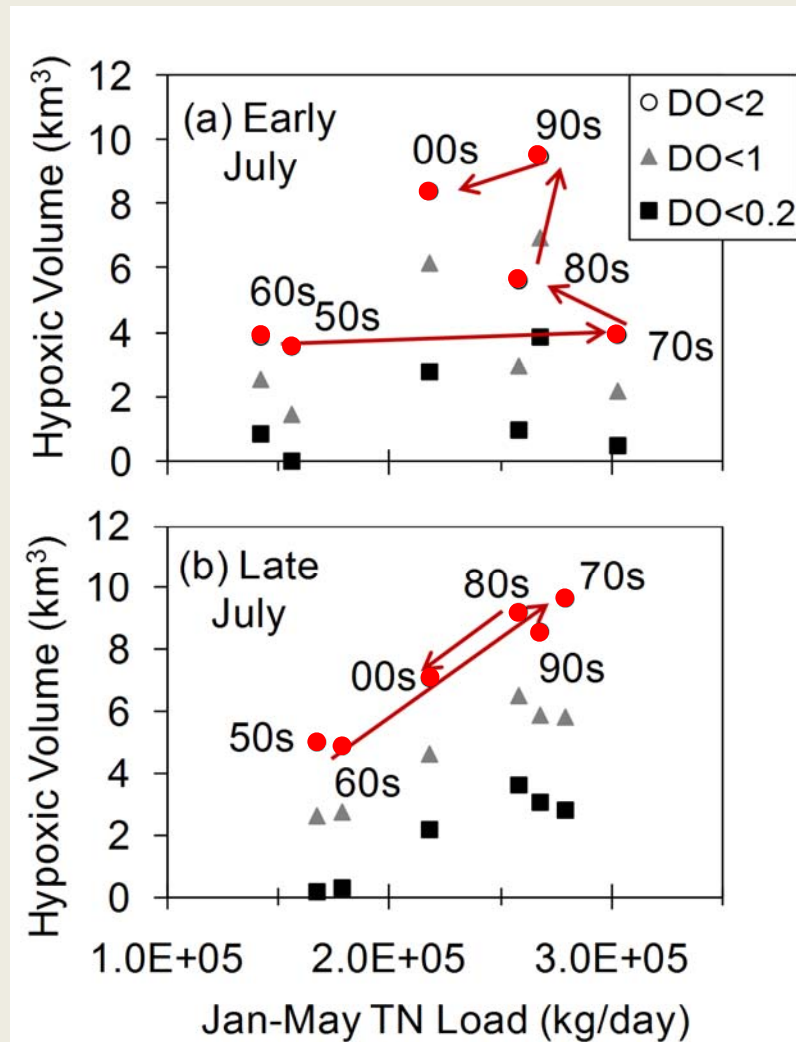
Murphy, Kemp, and Ball . 2011.  
*Estuaries and Coasts* 34:1293-1309.

# Hypoxia trends in relation to N-loading



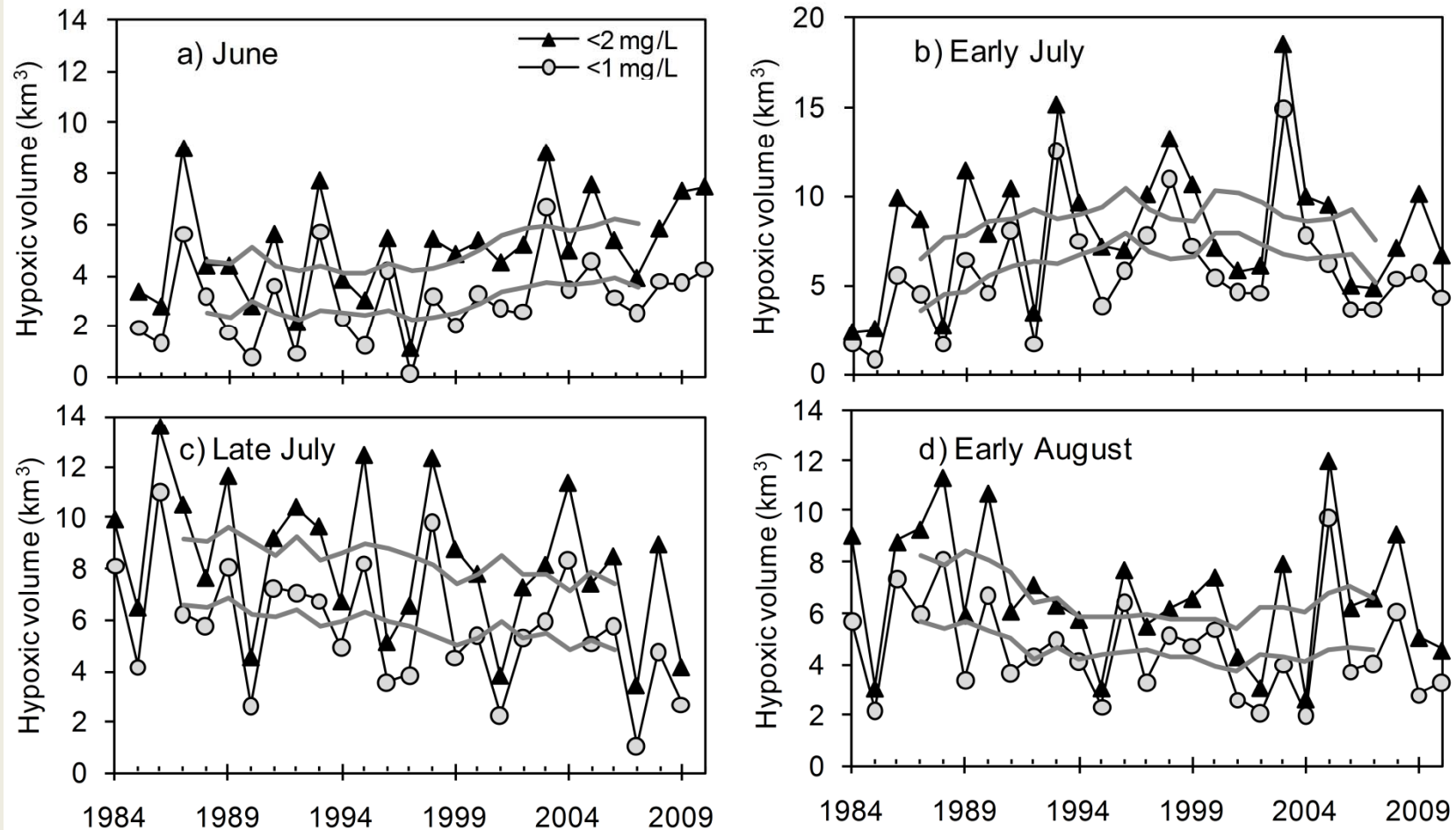
- Early July hypoxia shows increase since 1980
- Inter-annual variations blur long-term trends; clarify with running means
- Late July hypoxia has actually declined since 1980s
- N-loading increased until mid-1980s, then declined gradually into 2000s
- Hypoxia and late July N-load significantly correlated ( $p < 0.001$ )

## Decadal average hypoxic volume vs. TN



From: Murphy *et al.* 2011.

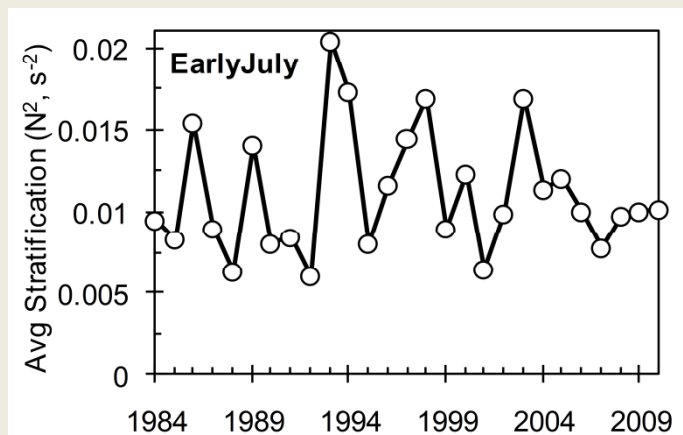
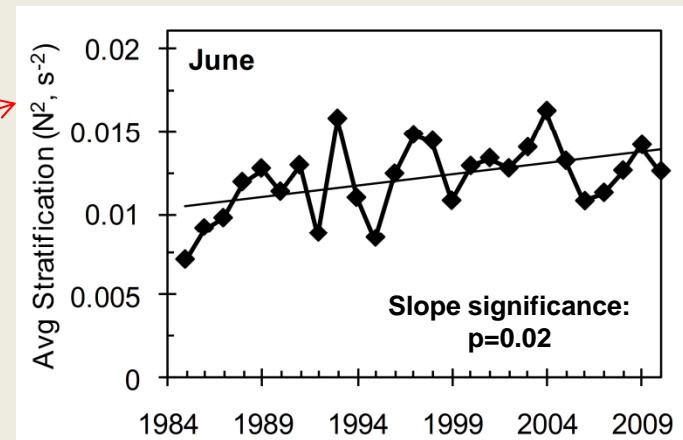
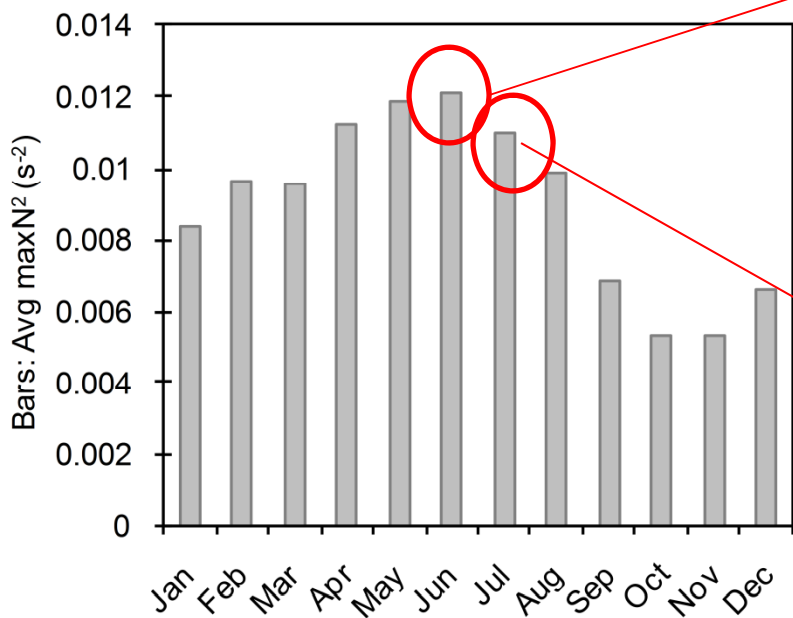
## Entire summer hypoxic volume trends



From: Murphy *et al.* 2011.

# Stratification long-term patterns

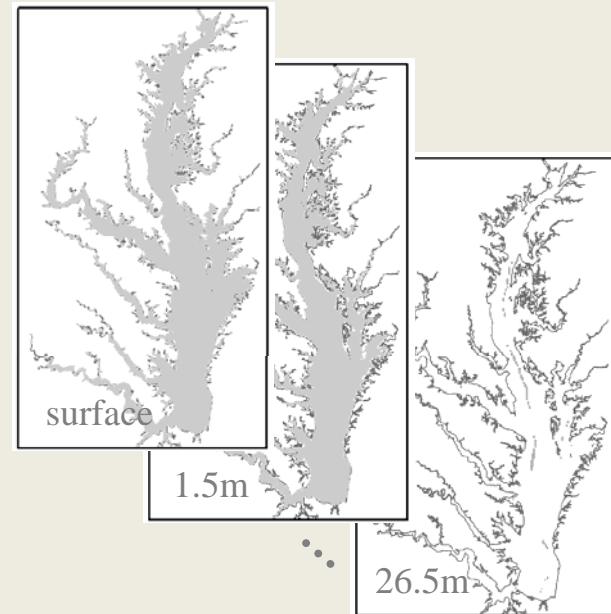
Average Monthly Main Channel Stratification Strength



# Outline

- What method(s) do you use for assessing DO?
- Why did you choose that methodology?
- What have you found drives DO patterns in the Bay?
  - Main channel hypoxia drives full-bay long-term trends
- What lessons have you learned that we should consider as we move forward?

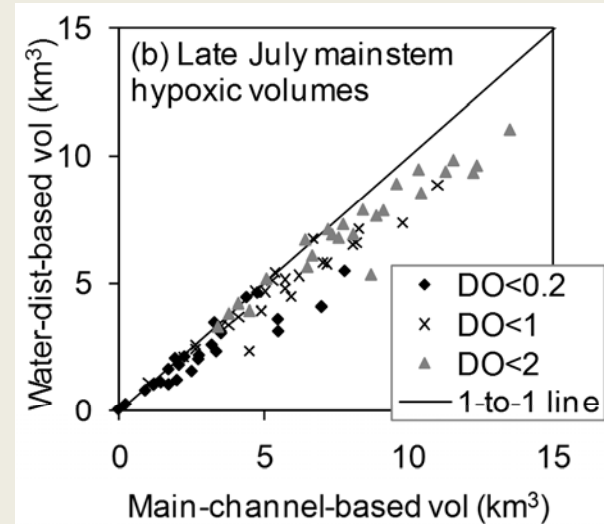
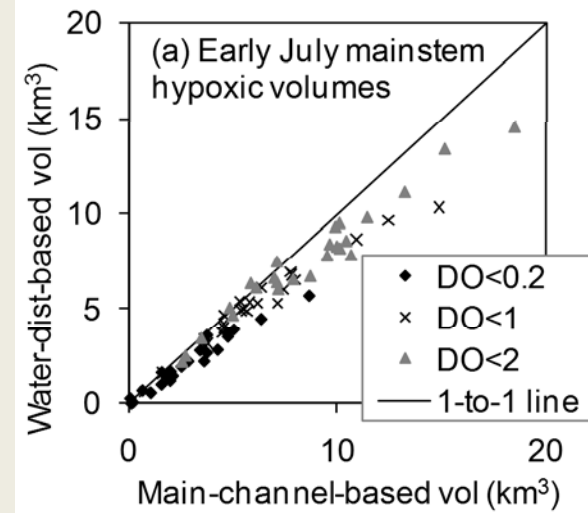
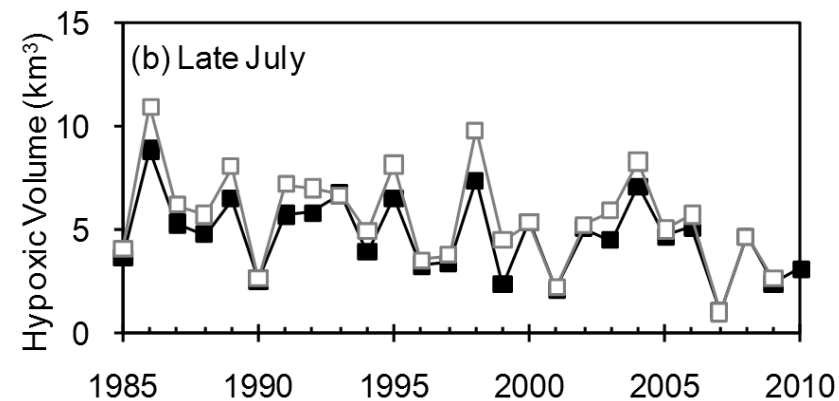
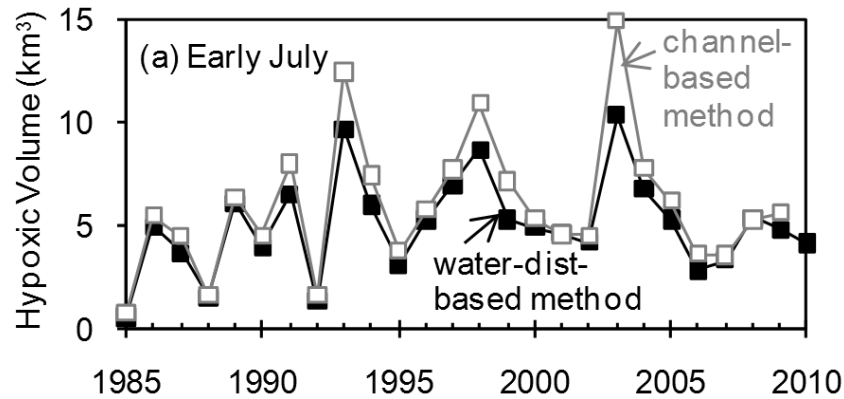
## *Full-bay method*





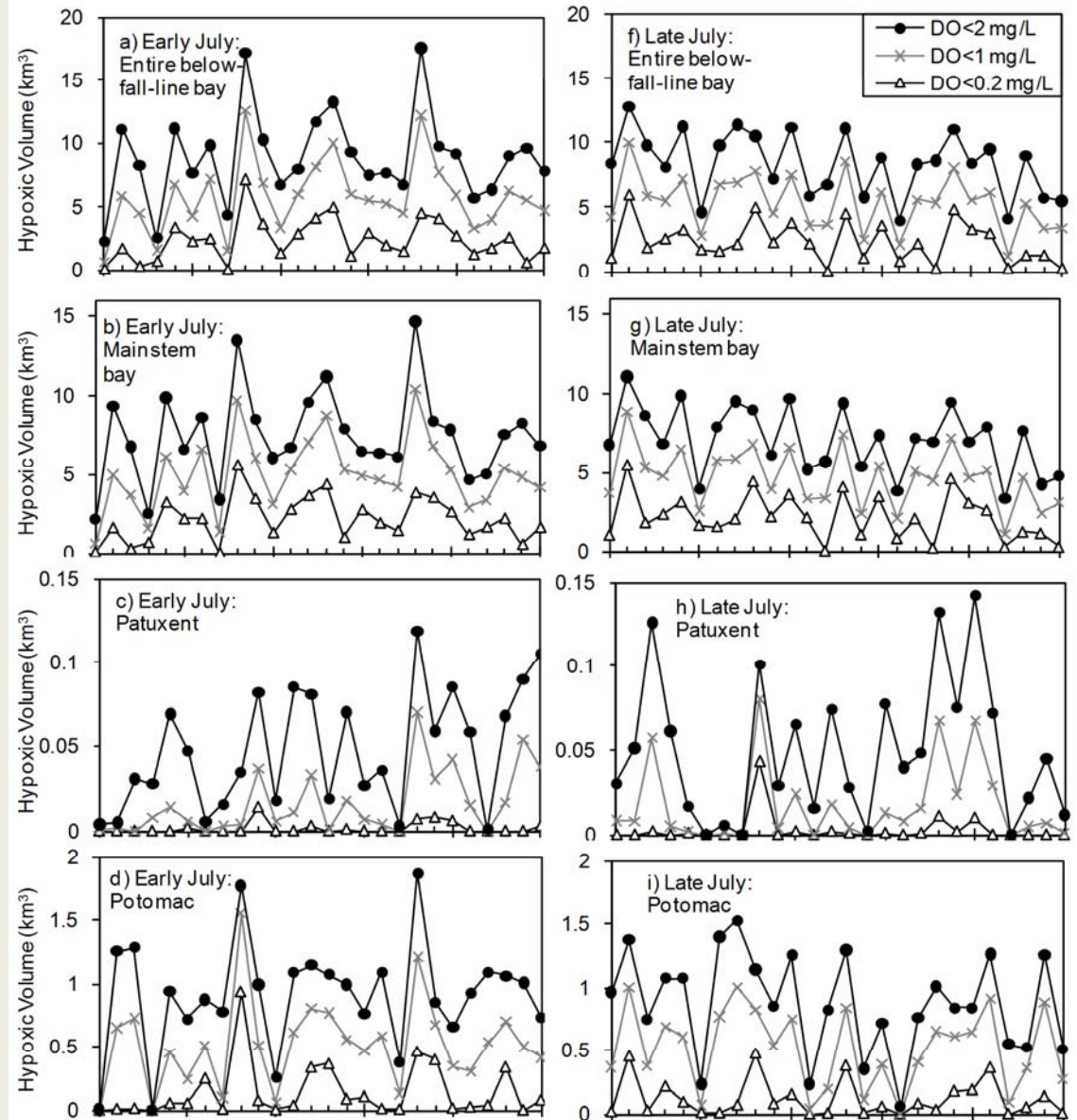
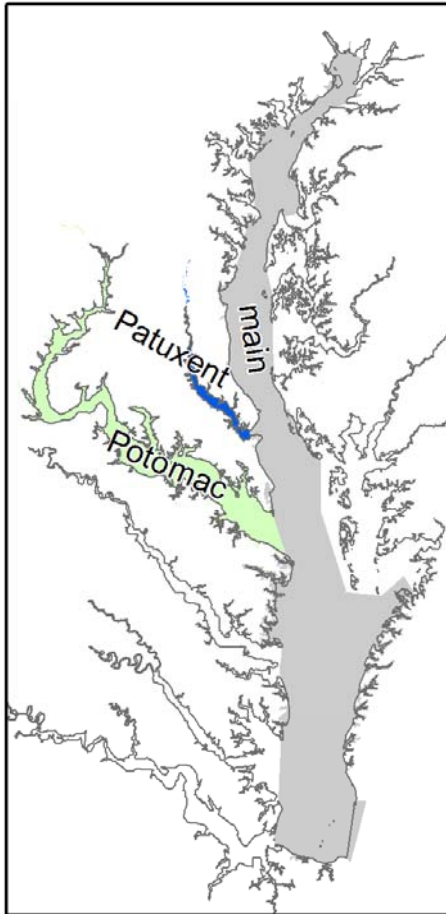
# Re-evaluation of hypoxic volume trends

Mainstem Hypoxic Volume ( $\text{DO} < 1 \text{ mg/L}$ )





## Results Part 2: Water distance



# Conclusions

- What have you found drives DO patterns in the Bay?
  - A trend in early summer stratification appears to have lead to an increase in early summer hypoxic volume.
  - The different early and late summer hypoxic volume trends observed are not artifacts of the main-channel-only interpolations.
  - More specific analysis could be done in tributaries.
- What lessons have you learned that we should consider as we move forward?
  - Timing is important when analyzing hypoxia.
  - Average monthly or summer hypoxia estimates may result in lost information.

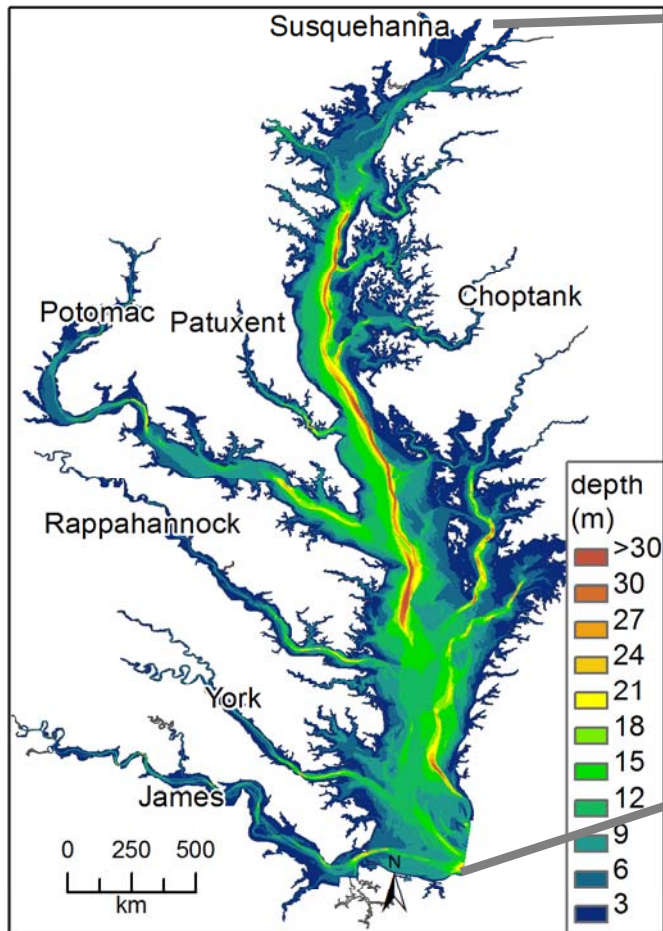
# Acknowledgements

- JHU Advisor: Bill Ball
- Committee members: Grace Brush, Frank Curriero, Dominic Di Toro, Michael Kemp
- “Distance” sources: Eric Perlman and Andrew Miglino
- CBEO team: Randal Burns, Damian Brady, Jen Bosch, Jeremy Testa, Maureen Brooks
- Data sources: CBP, CBI, USGS, NOAA, MDNR, VECOS
- NSF Grants no. 0618986 and 0854329
- EA Science, Engineering, Technology

Extras

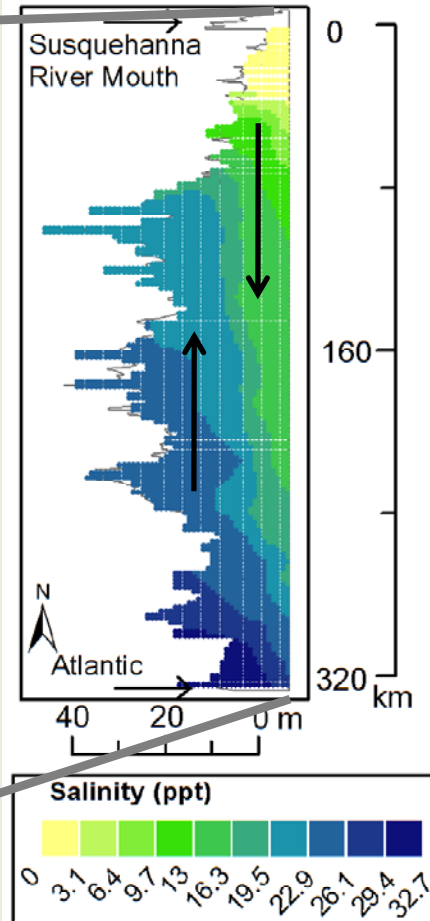
# Chesapeake Bay

Bay bathymetry



Bathymetry data from NOAA

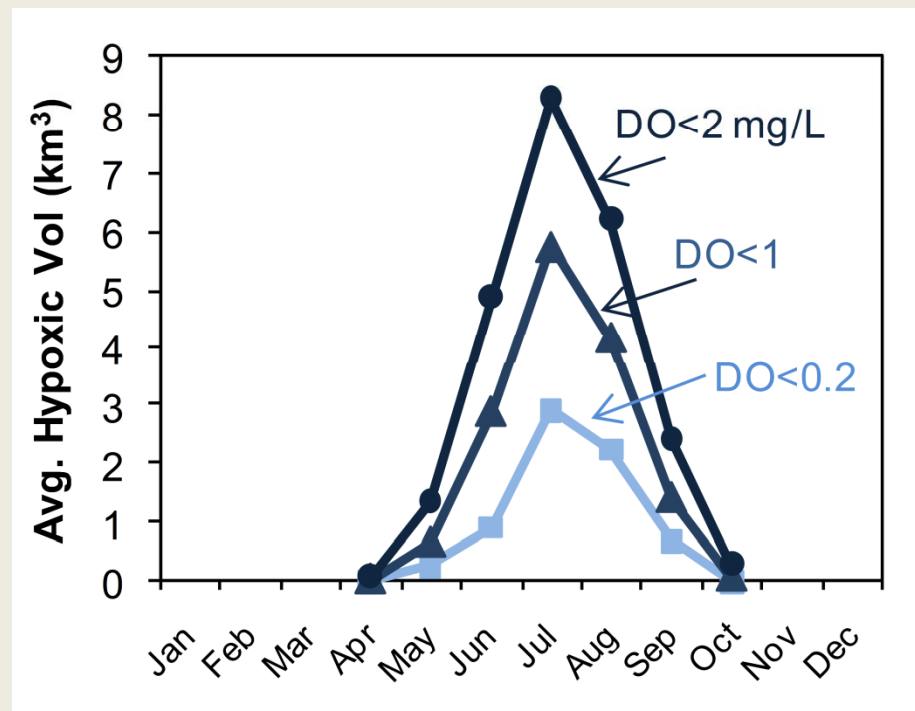
Example main channel salinity



- Bay is on average 6.5 m deep, but a much deeper channel follows the center of the Bay
- Most of the year, the Bay is stratified with fresher water on top, saltier water on bottom

## Hypoxic volume: Temporal extent

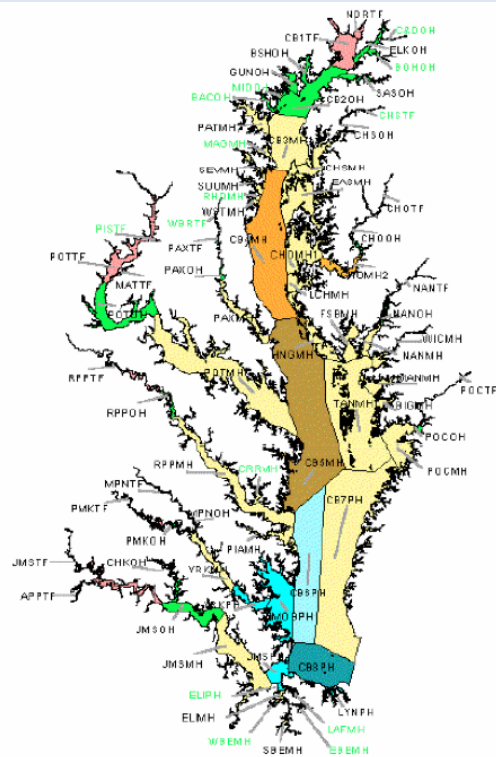
Hypoxic volume = volume of water with dissolved oxygen (DO) below certain critical values



Data summarized from: Murphy, Kemp, and Ball. 2011.  
*Estuaries and Coasts* 34:1293-1309.

# Interpolation method used for management

## 1. Split bay into segments and 1-m depth layers

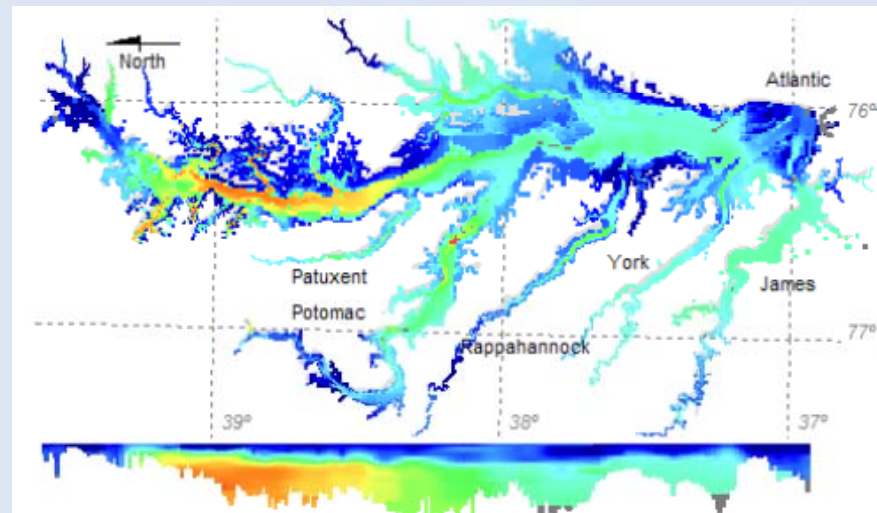


Chesapeake Bay Program 1998 segmentation design.

## 2. Interpolate samples in or near each segment using IDW

$$\hat{Y}(s_0) = \sum_{i=1}^n w(s_i) Y(s_i) \quad w(s_i) = \frac{1/d_i^2}{\sum_{j=1}^n (1/d_j^2)}$$

#### 4. Combine interpolation output by segment and layer





# Variogram

*Classical Variogram Estimator*

$$2\hat{\gamma}(\mathbf{h}) = \frac{1}{|N(\mathbf{h})|} \sum_{N(\mathbf{h})} (y(\mathbf{s}_i) - y(\mathbf{s}_j))^2$$

- $N(\mathbf{h})$  contain the pairs of locations  $(\mathbf{s}_i, \mathbf{s}_j)$  that are a distance  $\|\mathbf{h}\|$  apart,  $\mathbf{h} = \mathbf{s}_i - \mathbf{s}_j$ .
- $|N(\mathbf{h})|$  is the number of such pairs.

From Frank Curriero



# Some notes on kriging

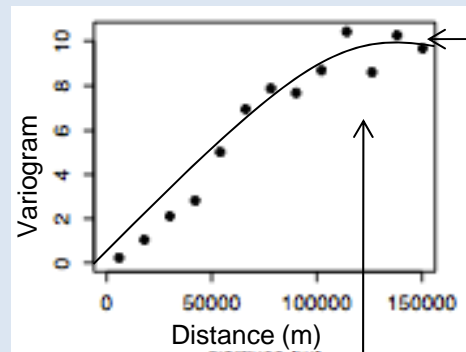
## Variogram

1. Examine an empirical variogram of observed data  $Y_2$  as a function of the Euclidean distance between points

$$2\hat{\gamma}(h) = \frac{1}{|N(h)|} \sum_{N(h)} [Y(s_i) - Y(s_j)]^2$$

$s_{i,j}$  are locations

$N(h)$  is number of pairs distance  $h$  apart



2. Fit a valid variogram model

## Estimate unknown values $Y_1$ given $Y_2$

3. Use variogram model to fill in covariance matrix:

$$\Sigma = \begin{bmatrix} \Sigma_{11} & \Sigma_{12} \\ \Sigma_{21} & \Sigma_{22} \end{bmatrix}$$

4. Use the conditional distribution of the unobserved data given the observed to generate estimates for  $Y_1$

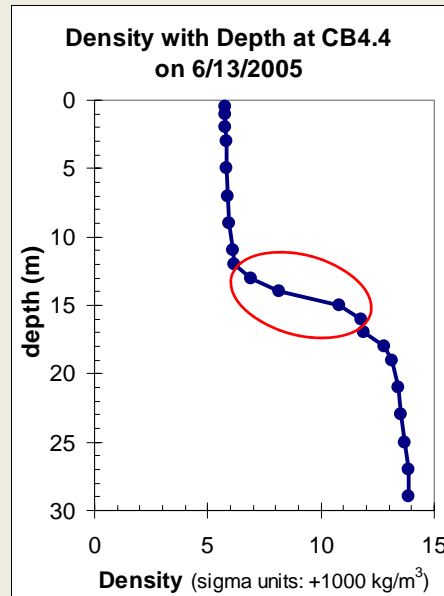
$$E(Y_1 | Y_2) = \mu_1 + \Sigma_{12} \Sigma_{22}^{-1} (Y_2 - \mu_2)$$

$\mu_2$  is mean of observed data,  $\mu_1$  is mean of unknown values (predicted using stationarity assumption)

Optional: Include covariates to account for some spatial variability in  $\mu$  (universal kriging)

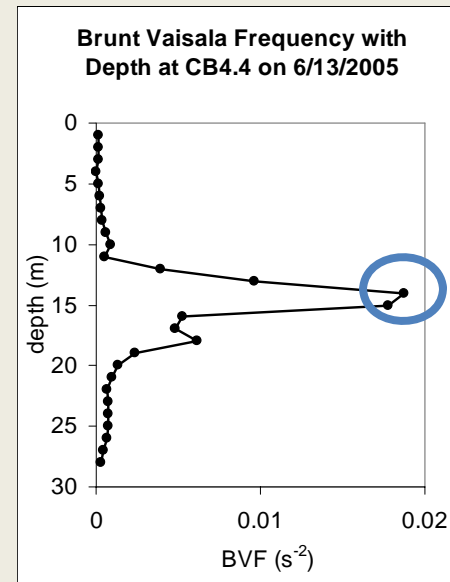
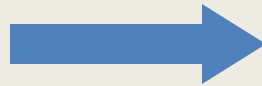
# Hypoxia/Stratification

# Calculation of stratification strength

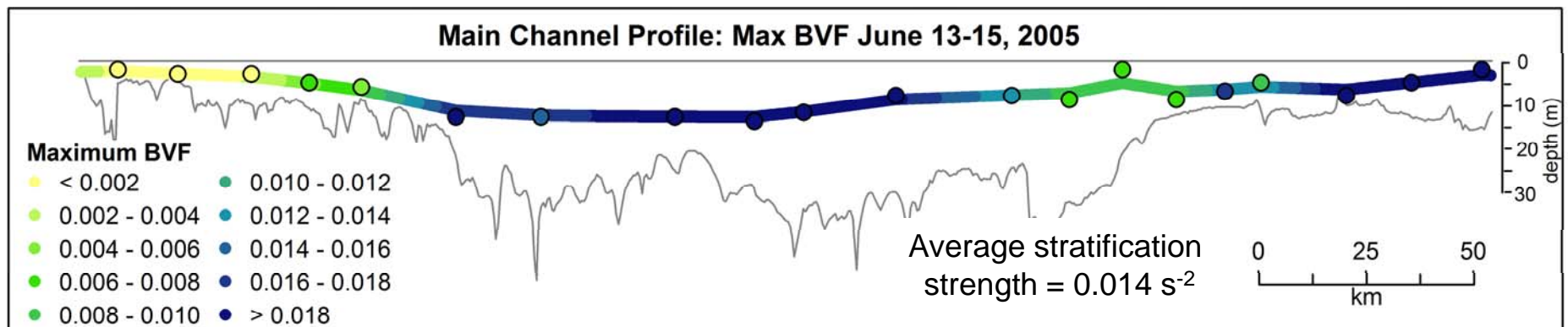
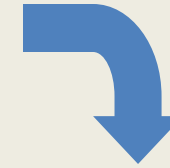


1. Calculate Brunt Väisälä Frequency

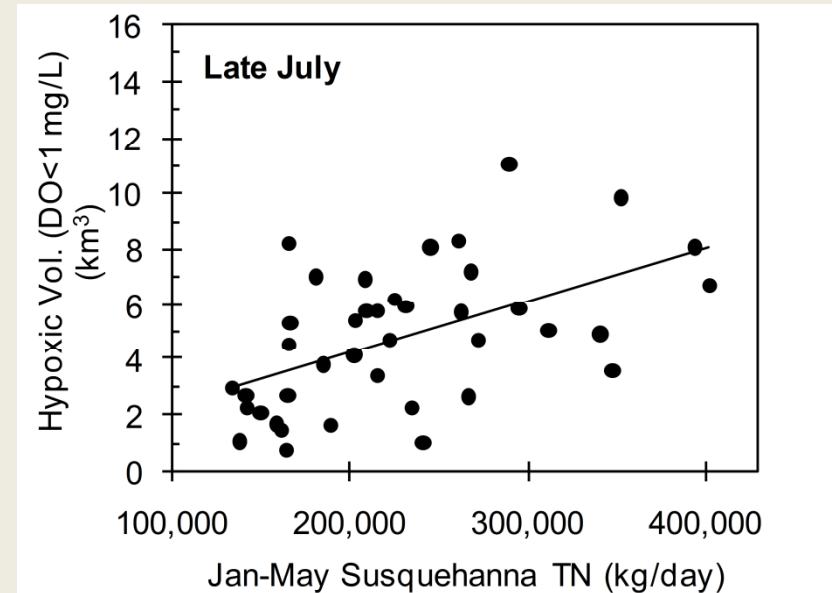
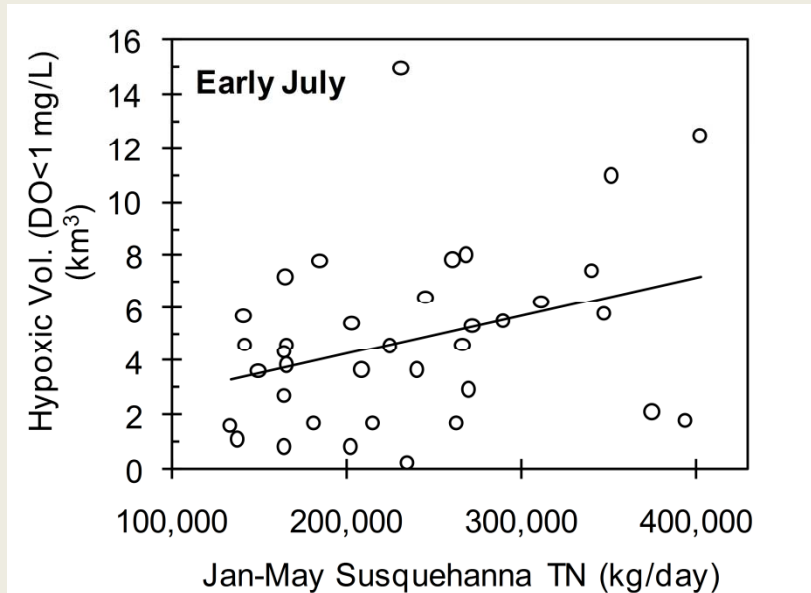
$$BVF = \frac{g}{\rho} \frac{\partial \rho}{\partial z}$$



2. Interpolate maximum Brunt Väisälä Frequency (BVF)

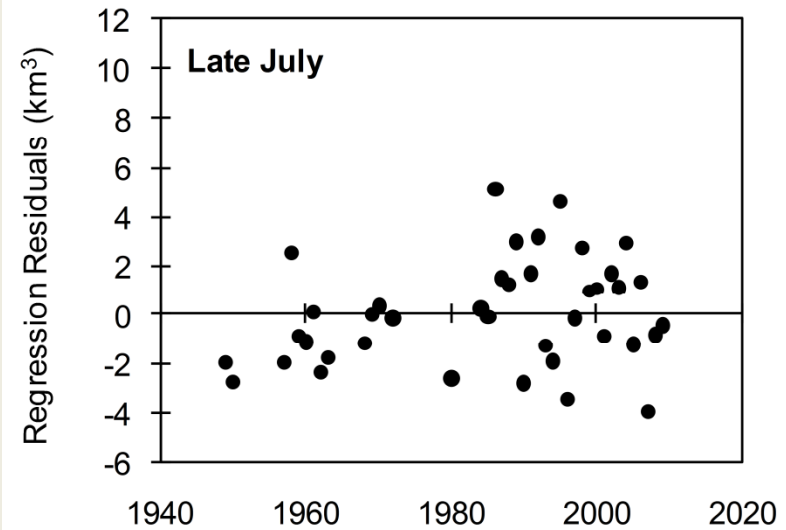
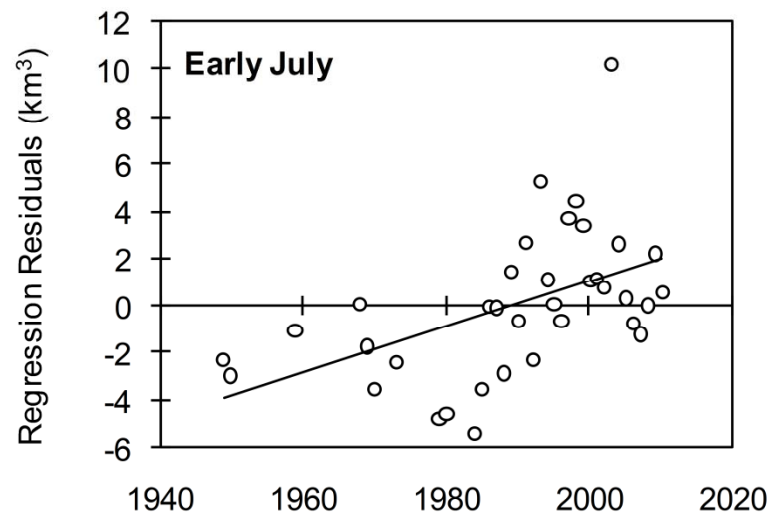


# Hypoxic volume vs. TN



Regression Model	HypoxicVol= $\beta_0 + \beta_1(TN)$
Early July (DO<1 mg/L)	$R^2=0.12$ , p-value=0.04
Late July (DO<1 mg/L)	$R^2=0.28$ , p-value<0.001

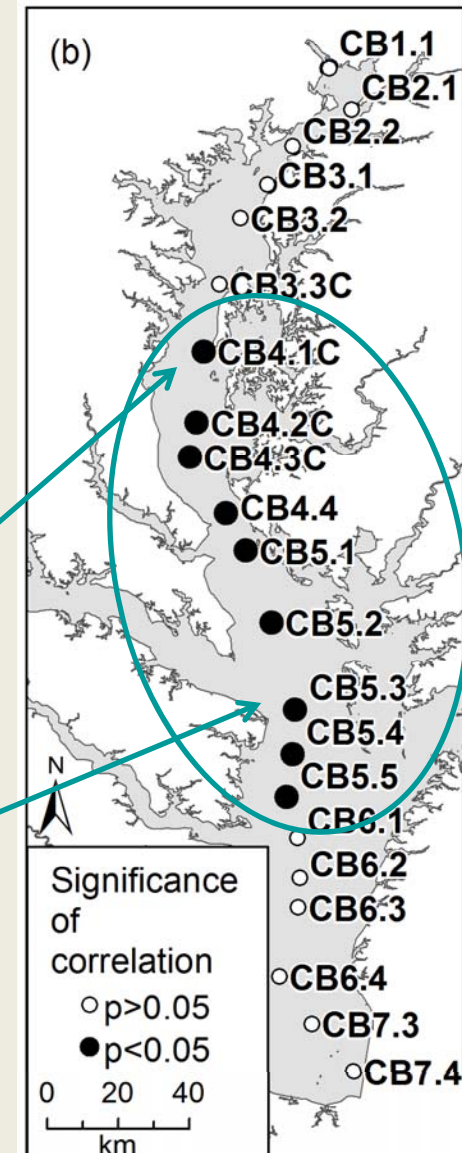
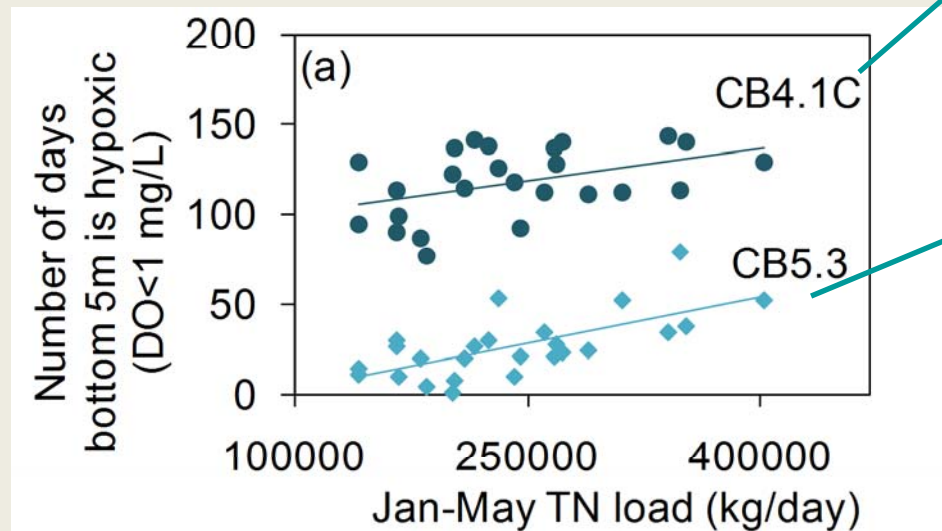
# Hypoxic volume vs. TN residuals



Regression Model	HypoxicVol= $\beta_0 + \beta_1(\text{TN})$	Residuals= $\beta_0 + \beta_1(\text{year})$
Early July (DO<1 mg/L)	$R^2=0.12$ , p-value=0.04	$R^2=0.25$ , p-value=0.001
Late July (DO<1 mg/L)	$R^2=0.28$ , p-value<0.001	$R^2=0.07$ , p-value=0.1

# Hypoxia Duration

- Number of days that bottom 5 m of water are hypoxic
- In mid-Bay, the hypoxia duration is significantly correlated with Jan-May TN loads



From: Murphy *et al.* 2011.

## Regression for hypoxic volume (DO<1 mg/L)

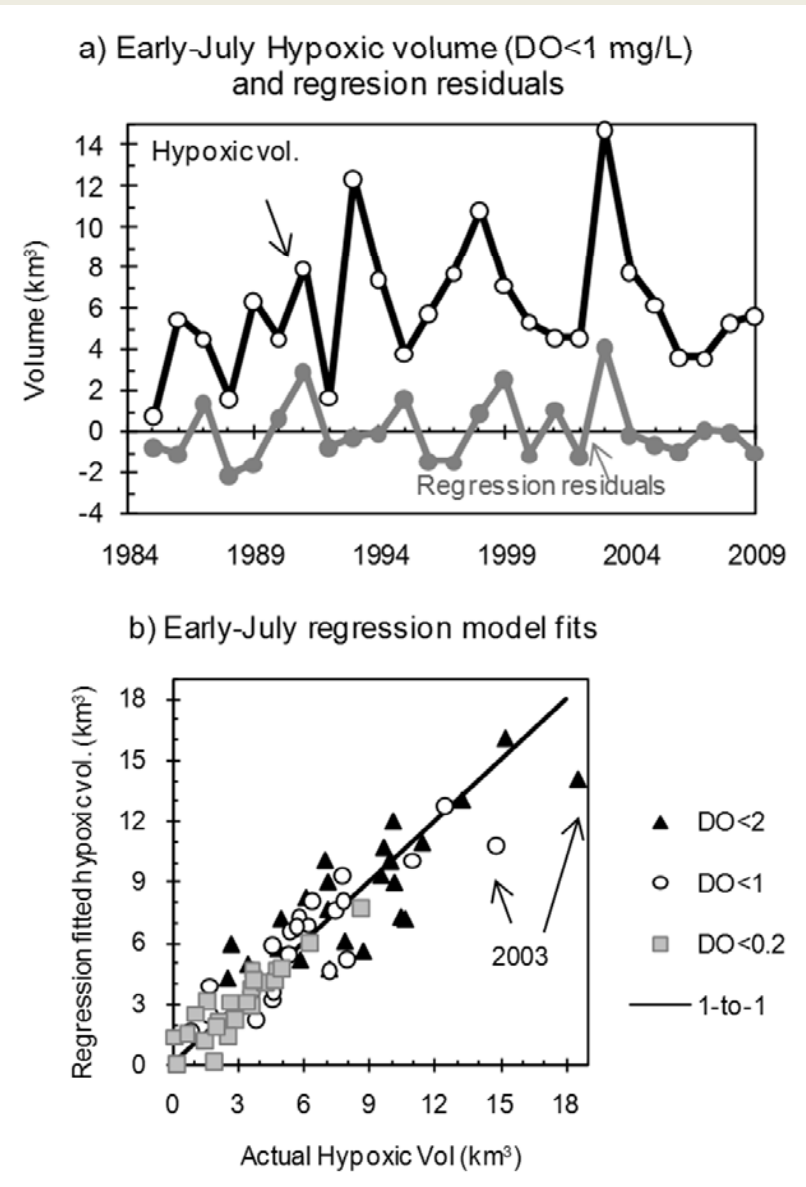
$$\text{HypoxicVol}_{\text{EarlyJuly}} = \beta_0 + 1.2 (\text{Stratification}_{\text{June}}) + 1.8 (\text{Stratification}_{\text{EarlyJuly}}) + 0.5 (\text{VolBelowPyc}_{\text{EarlyJuly}}) + \varepsilon$$

$$R^2=0.77, p<0.001$$

*June stratification is only variable that explains the temporal increase*

From 1985-2009

## Results Part 1: Main-channel hypoxia



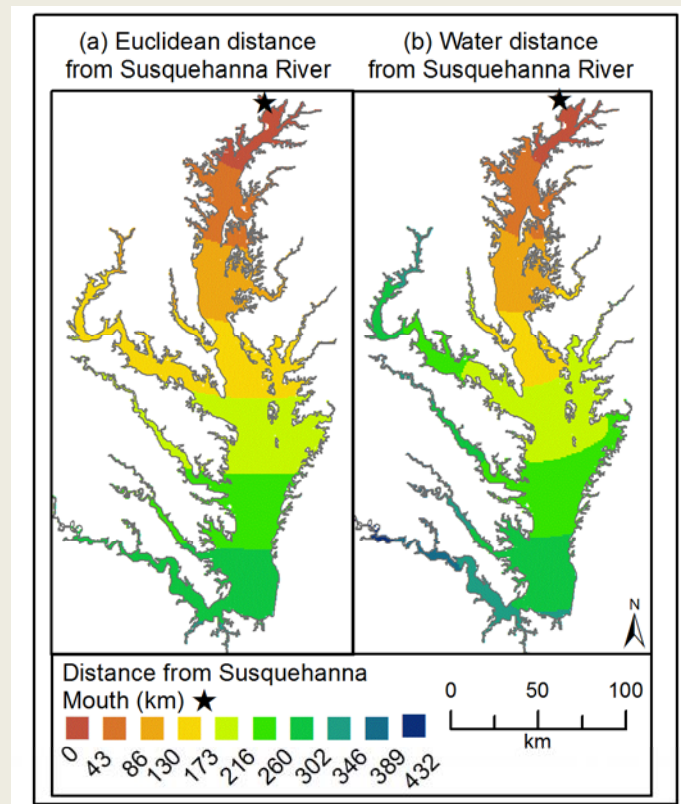
From: Murphy *et al.* 2011.

Water distance



## Water distance

- Distance between locations through water
- Computed between 11,534 grid and sample locations



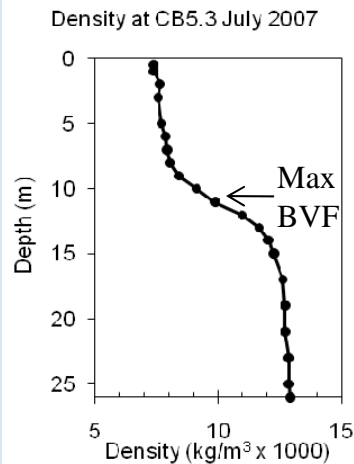
Water distances computed by  
Eric Perlman (was at JHU CS)

# Water-distance-based method: Layers

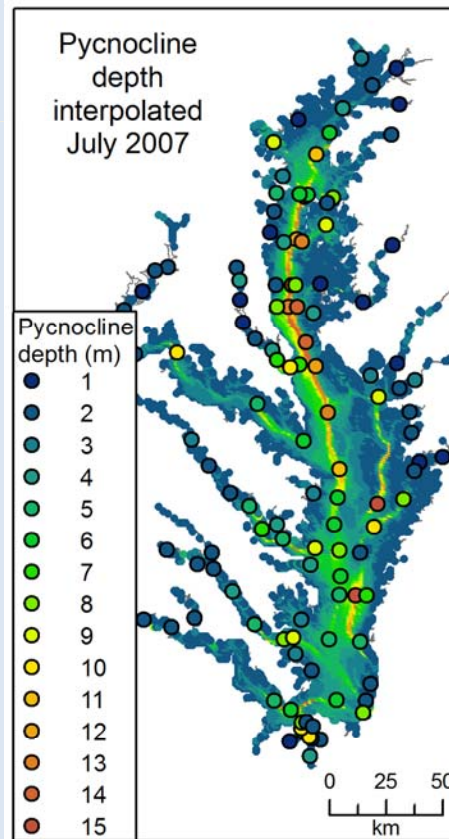
## Method 2: Full Bay

### 1. Identify depth of pycnocline

$$BVF = \frac{g}{\rho} \frac{\partial \rho}{\partial z}$$



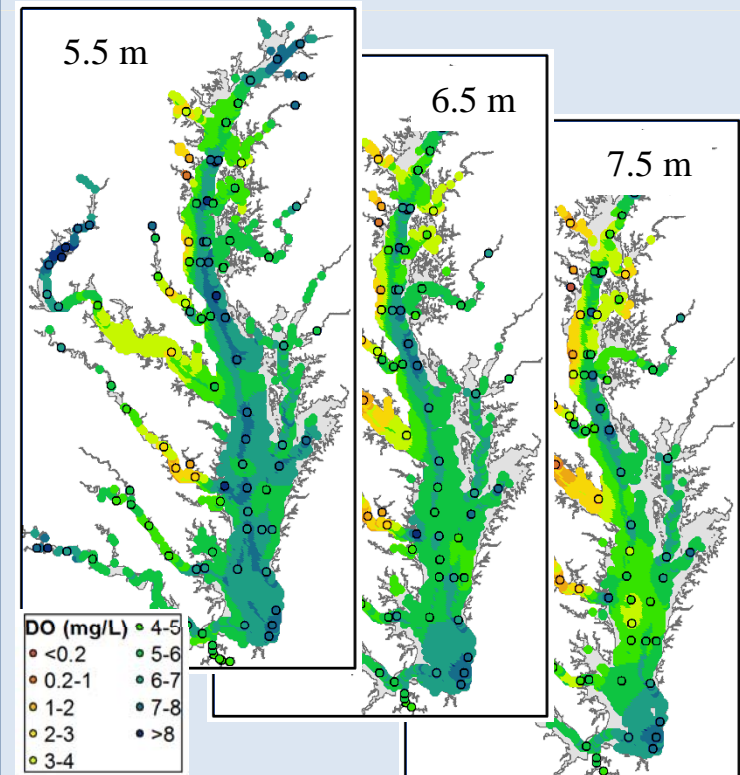
### 2. Interpolate throughout bay



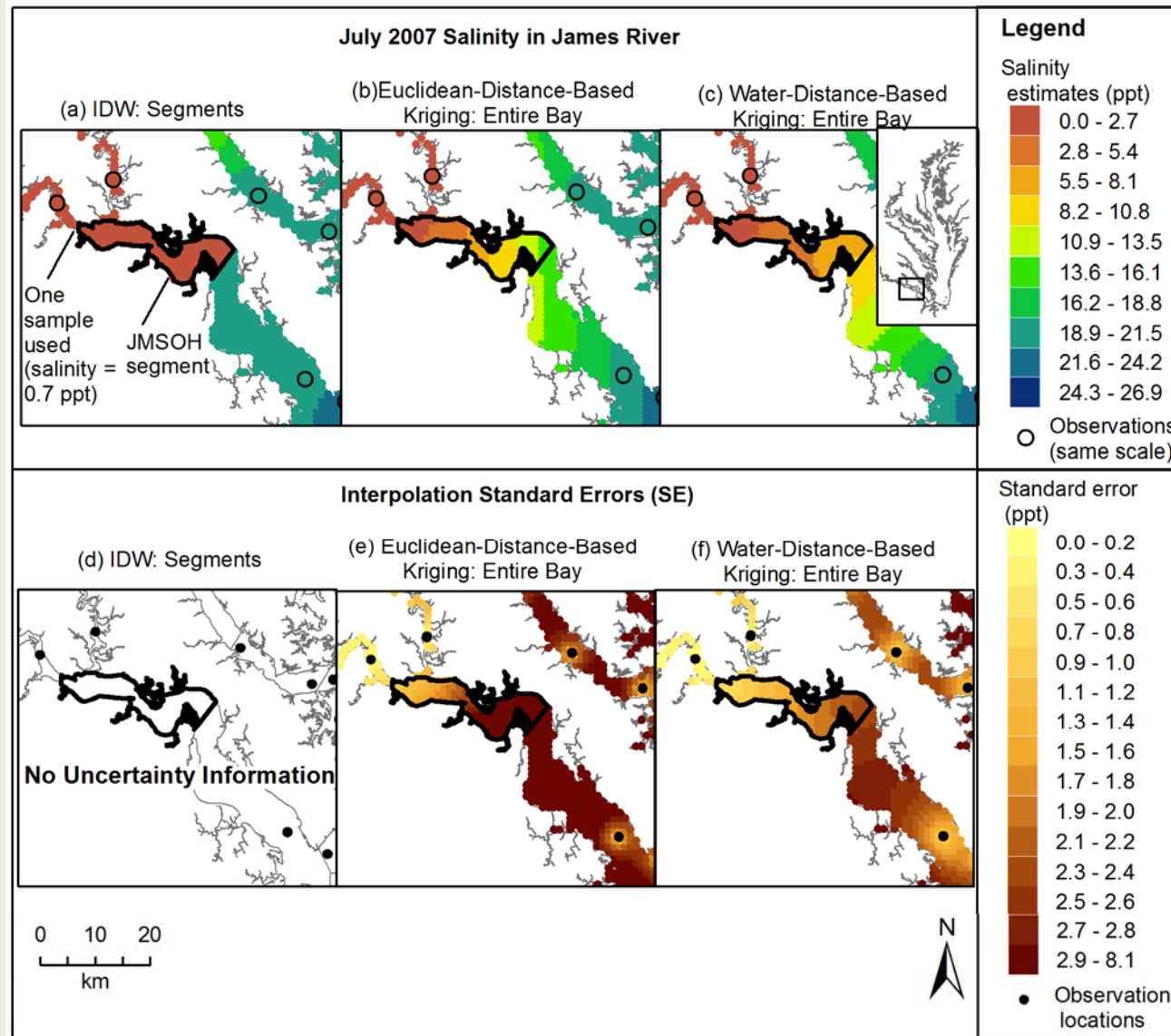
3. Use pycnocline depth indicator (P) as a covariate to interpolate DO at location s

$$\text{DO}(s) = \beta_0 + \beta_1 P(s) + \varepsilon(s)$$

### 4. “Stack” interpolated DO layers to generate 3-D estimate



# Interpolation Results



# Cross Validation Summary

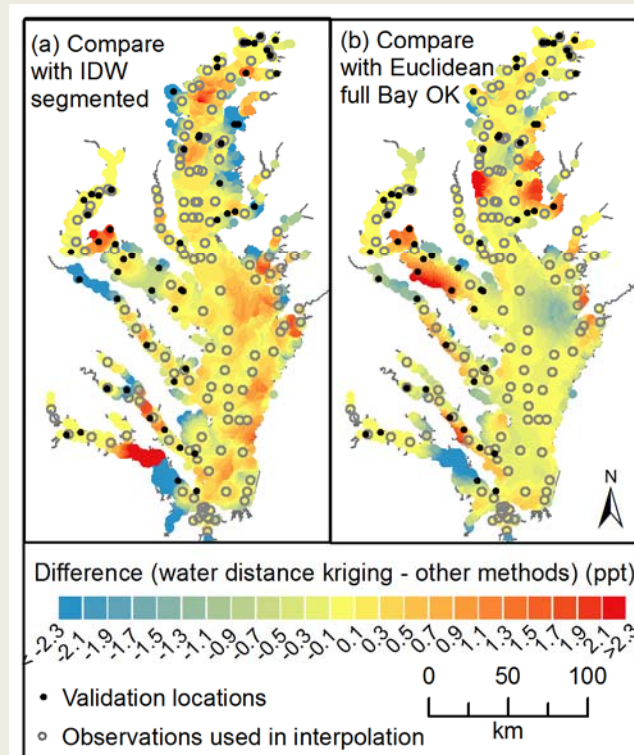
**RMSE for July 2007 (and percent of time method has lowest RMSE for all 2007/2008 sets)**

Parameter	IDW in segments	Kriging with Euclidean distance	Kriging with water distance
Water temp	0.79 (12%)	0.76 (8%)	<b>0.74 (80%)</b>
Salinity	2.3 (0%)	1.5 (0%)	<b>1.1 (100%)</b>
DO	1.5 (0%)	<b>1.2 (8%)</b>	<b>1.2 (92%)</b>
NH <sub>4</sub>	0.023 (16%)	<b>0.022 (20%)</b>	0.023 (64%)
NO <sub>2</sub> + NO <sub>3</sub>	0.18 (12%)	0.25 (16%)	<b>0.17 (72%)</b>
PO <sub>4</sub>	0.0079 (4%)	<b>0.0073 (12%)</b>	0.0075 (84%)
Chl-a	<b>15 (8%)</b>	<b>15 (32%)</b>	<b>15 (60%)</b>

## Validation Test

- Identify continuous monitoring stations from Maryland Department of Natural Resources and Virginia Estuarine and Coastal Observing System
- Average continuous observations 1 hour before and after the time of closest CBP sample
- Compare validation observation to interpolation output

Surface Salinity July 2007: Difference between water distance kriging and other methods and validation locations



Validation: <http://www3.vims.edu/vecos/>  
<http://mddnr.chesapeakebay.net/newmontech/contmon/index.cfm>

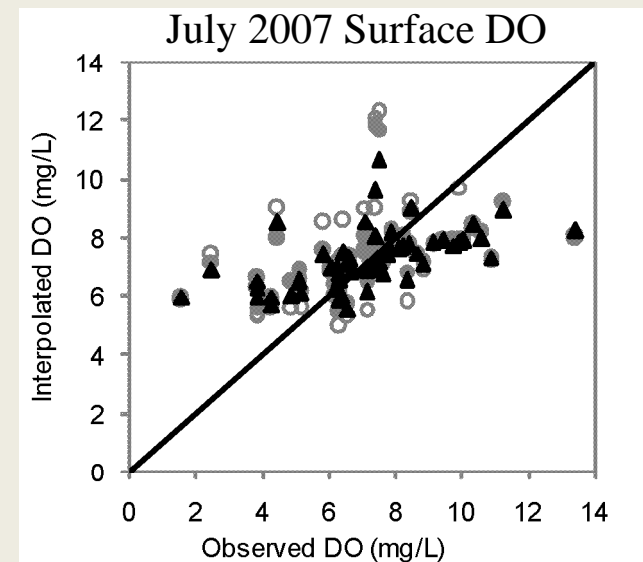
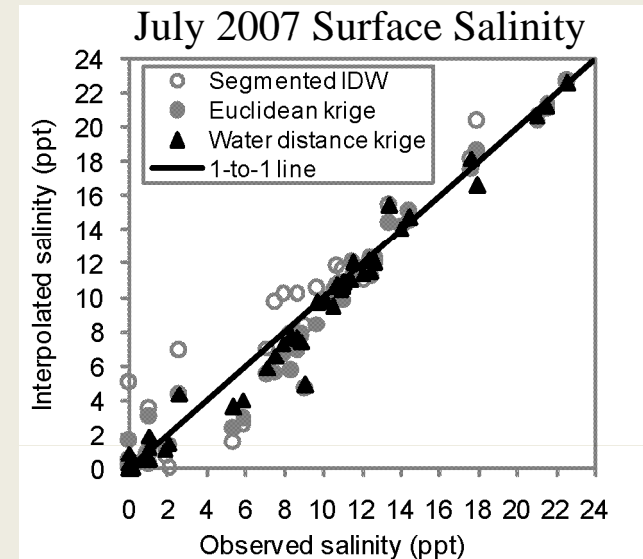
# Validation Results

## Root Mean Squared Errors of Estimates

	IDW in segments	Kriging with Euclidean distance	Kriging with water distance
Water temp	1.0	0.88	<b>0.84*</b>
Salinity	1.4	1.2	<b>0.92</b>
DO	2.0	1.9	<b>1.8</b>
Chl-a	29	24	<b>21</b>

\*Uses a covariate, depth

**Validation sets:** Maryland Department of Natural Resources and Virginia Estuarine and Coastal Observing System continuous monitoring data





# Hypoxic volumes computed using water-distance-based method

*Results Part 2: Water distance*

